



IUSS  
FERRARA 1391



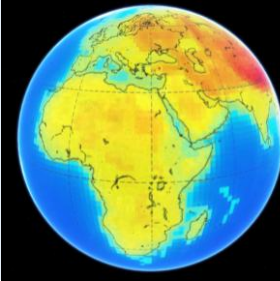
# A refined reference Earth model for the geo-neutrino studies at Borexino

Viacheslav Chubakov

TUTORI: Prof. Giovanni Fiorentini, PhD. Fabio Mantovani

Università di Ferrara

# Summary



## Introduction:

- geoneutrino properties
- why do we need a refined reference Earth model (RRM)

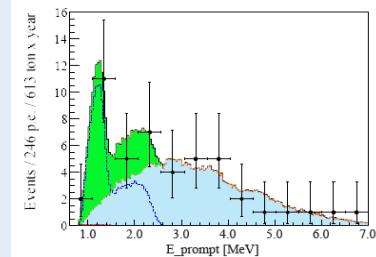
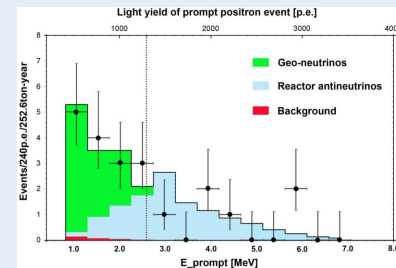
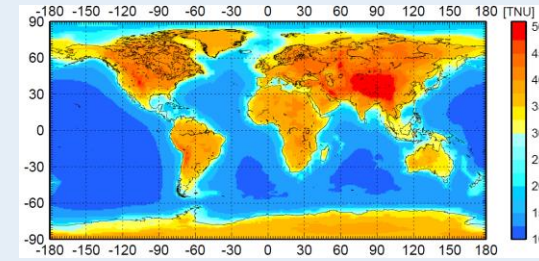
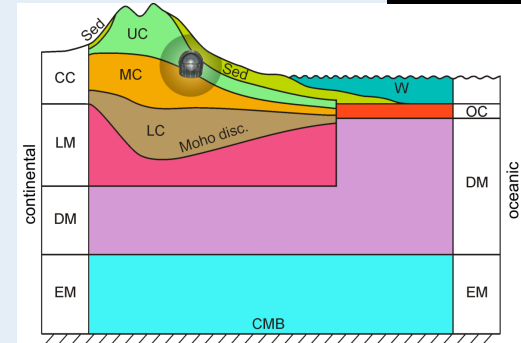
## Main ingredients of RRM

- Global model of crust
- Model of mantle
- Local model of crust around LNGS

## Antineutrino background from reactors

## Geoneutrino measurements of Borexino

## Conclusions



# Geo-neutrinos: anti-neutrinos from the Earth

$^{238}\text{U}$ ,  $^{232}\text{Th}$  and  $^{40}\text{K}$  in the Earth release heat together with anti-neutrinos, in a **well fixed ratio**:

Decay	$T_{1/2}$ [ $10^9$ yr]	$E_{\text{max}}$ [MeV]	$Q$ [MeV]	$\varepsilon_{\bar{\nu}}$ [ $\text{kg}^{-1}\text{s}^{-1}$ ]	$\varepsilon_H$ [W/kg]
$^{238}\text{U} \rightarrow ^{206}\text{Pb} + 8\ ^4\text{He} + 6e + 6\bar{\nu}$	4.47	3.26	51.7	$7.46 \times 10^7$	$0.95 \times 10^{-4}$
$^{232}\text{Th} \rightarrow ^{208}\text{Pb} + 6\ ^4\text{He} + 4e + 4\bar{\nu}$	14.0	2.25	42.7	$1.62 \times 10^7$	$0.27 \times 10^{-4}$
$^{40}\text{K} \rightarrow ^{40}\text{Ca} + e + \bar{\nu}$ (89%)	1.28	1.311	1.311	$2.32 \times 10^8$	$0.22 \times 10^{-4}$

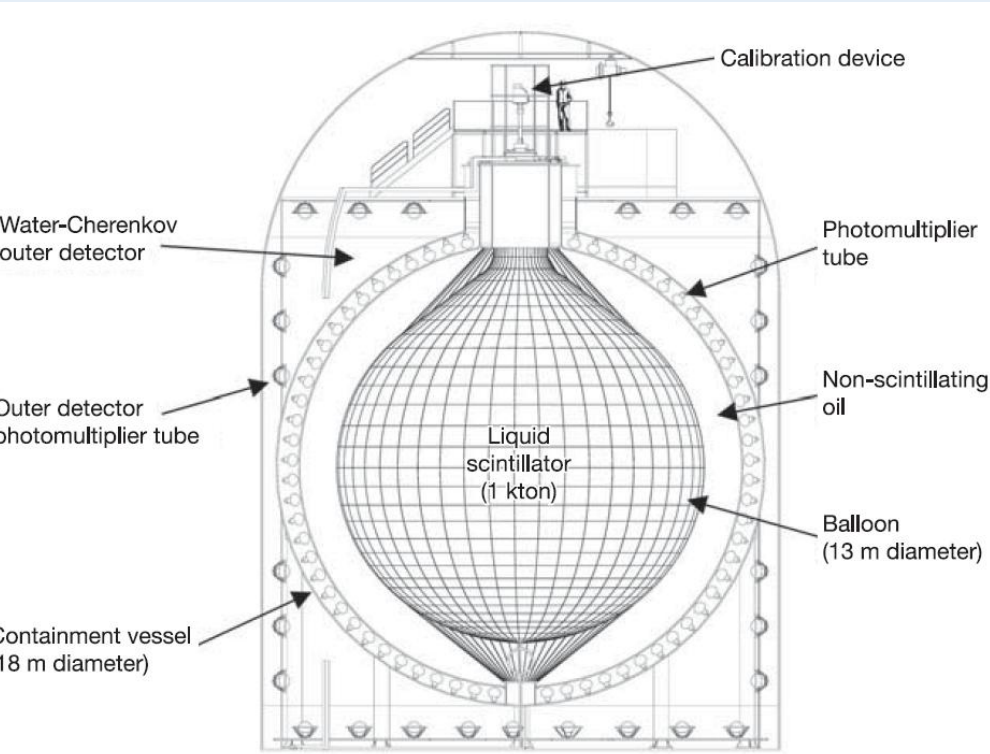
- Earth emits (mainly) antineutrinos  $\Phi_{\bar{\nu}} \sim 10^6 \text{ cm}^{-2}\text{s}^{-1}$  whereas Sun shines in neutrinos.
- A fraction of geoneutrinos from U and Th (not from  $^{40}\text{K}$ ) are above threshold for inverse  $\beta$  on protons:  $\bar{\nu} + p \rightarrow e^+ + n - 1.8 \text{ MeV}$
- Different components can be distinguished due to different energy spectra: e. g. anti-n with highest energy are from Uranium.
- Signal unit: **1 TNU** = one event per  $10^{32}$  free protons per year.

# Running experiments: KamLAND and Borexino



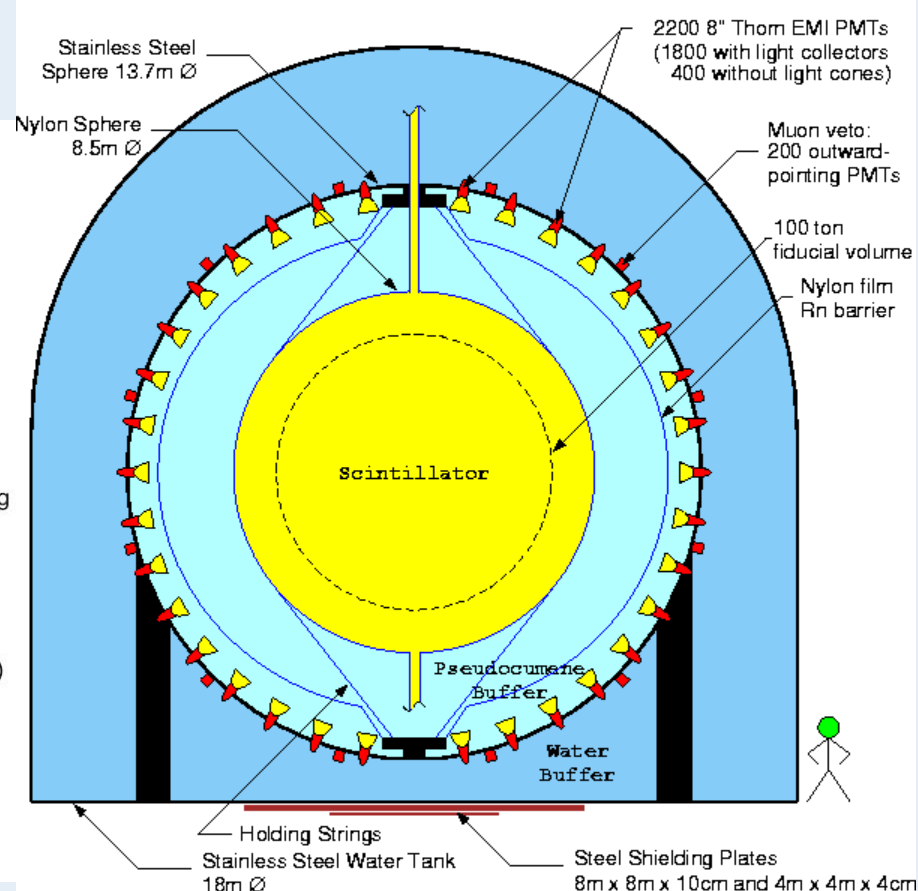
**Kamioka**  
**Liquid**  
**Scintillator**  
**ANTineutrino**  
**Detector**

1 kton LS, surrounded by 1845 PMT's



**Borexino**

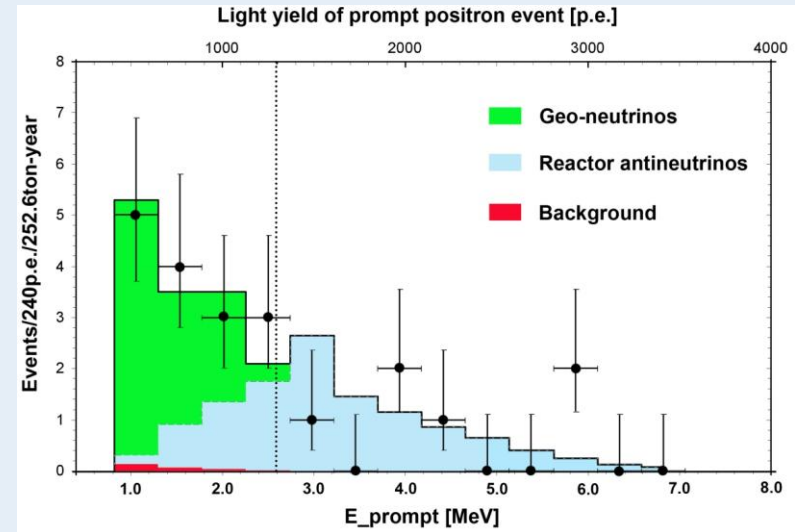
placed at Gran Sasso  
 National Laboratories  
 300 ton LS, surrounded  
 by 2200 PMT's



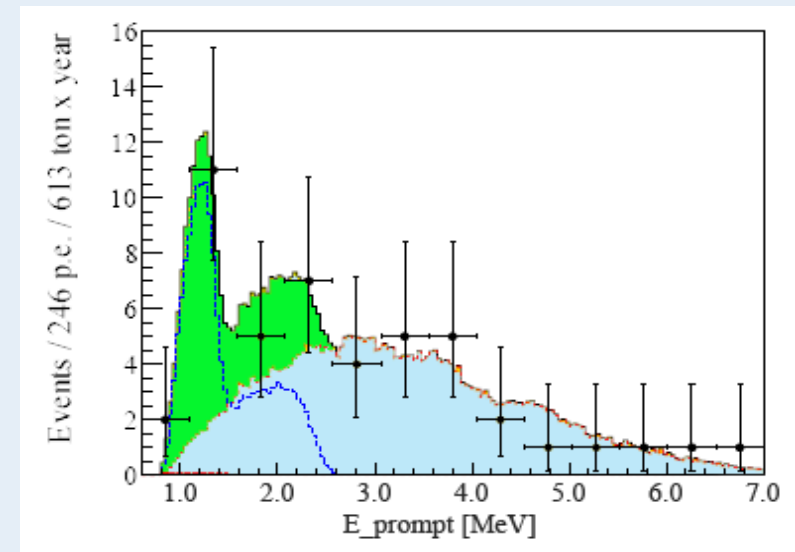
# Geo-neutrino measurements by Borexino



Period	Dec. 07 – Dec. 09	
Exposure p.yr	$0.15 \times 10^{32}$	
	events	TNU
Total (full sp.)	21	
Reactors	$10.7_{-3.4}^{+4.3}$	$64_{-22}^{+26}$
Geo-n	$9.9_{-3.4}^{+4.1}$	$69_{-22}^{+28}$
Background	$0.4_{-0.05}^{+0.05}$	



Period	Dec. 02 – Aug. 12	
Exposure p.yr.	$0.37 \times 10^{32}$	
	events	TNU
Total (full sp.)	46	
Reactors	$31.2_{-6.1}^{+7.0}$	$84.5_{-16.9}^{+19.3}$
Geo-nu	$14.3 \pm 4.4$	$38.8 \pm 12$
Background	$0.70 \pm 0.18$	

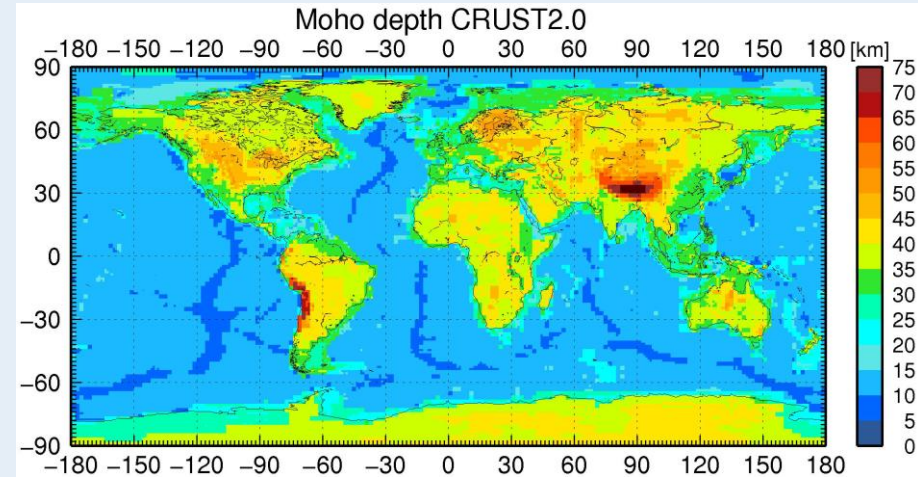


# Previous reference Earth models (RM)

Crust and mantle were treated as two separate reservoirs. In particular, the mantle was conventionally described as a spherically uniform shell between the crust and the core.

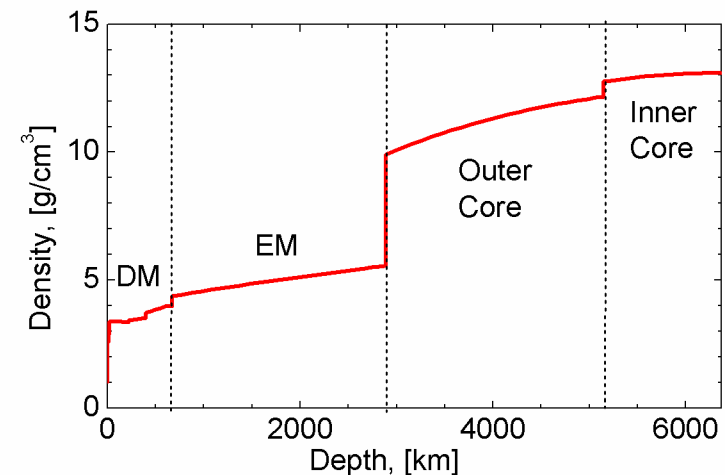
## 1) Crust:

- a) geophysical properties (thickness, density, etc.) were based on  $2^{\circ} \times 2^{\circ}$  CRUST2.0 map;
- b) composition was adopted from the world wide average abundances.

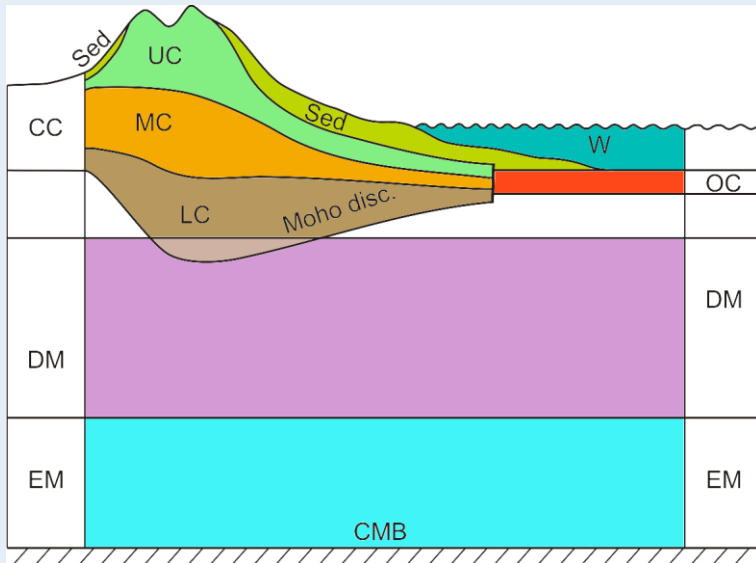


## 2) Mantle structure was based on PREM (Preliminary Earth Reference Model, a 1-D seismologically based global model)

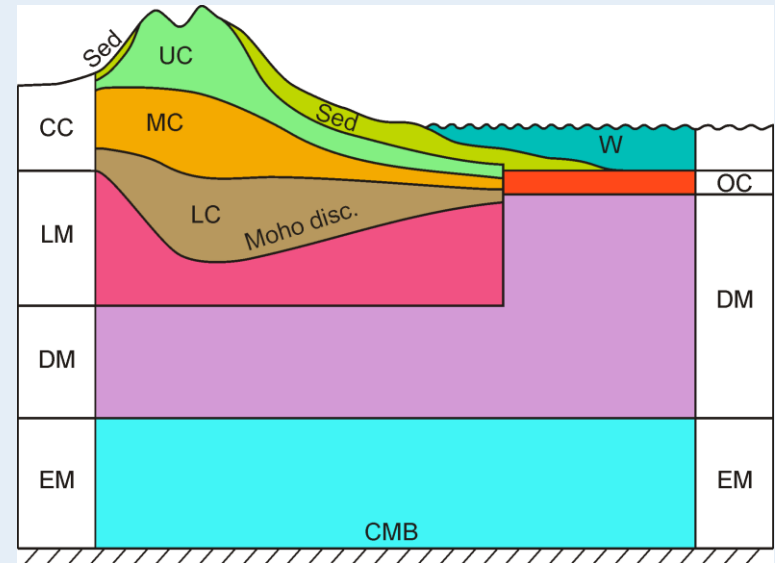
## 3) Core is considered to have negligible amounts of K, Th and U.



## Structure of RM



## Structure of RRM



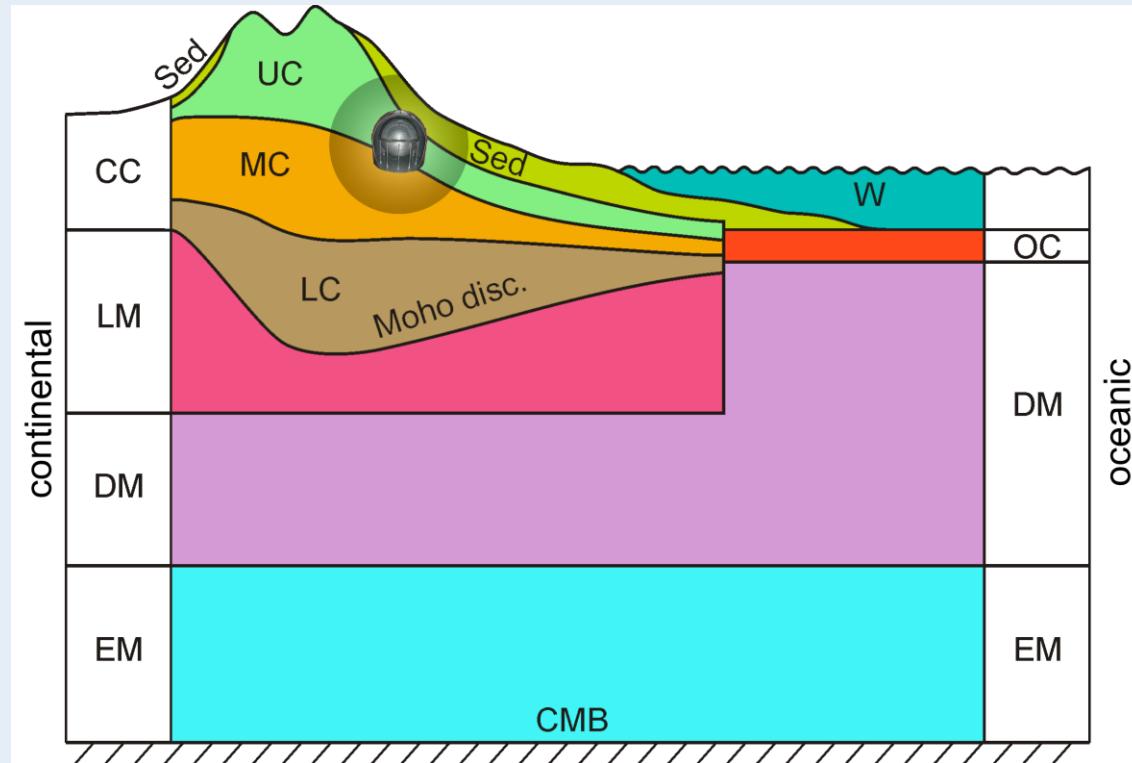
## Geoneutrino signal at LNGS predicted by different RM

Reference	$S_{\text{Gran Sasso}}$ TNU
Mantovani et al. 2004	$40.5 \pm 6.5$
Fogli et al., 2006	$40.5 \pm 2.9$
Enomoto et al. 2007	43.1
Dye, 2010	$42.0 \pm 7.2$

# Towards a Refined Reference Earth Model (RRM)\*

## The main updates of the RRM

- An updated  $1^\circ \times 1^\circ$  sedimentary layer
- Combination of 3 models for crust thickness
- New compilations of HPEs abundances in OC, Sed, UC
- New approach in the evaluation of U and Th abundances (and their uncertainties) in MC and LC based on seismic arguments
- Heterogeneous topography of crust bottom
- Introduction of CLM
- Propagation of uncertainties by using Monte Carlo simulation

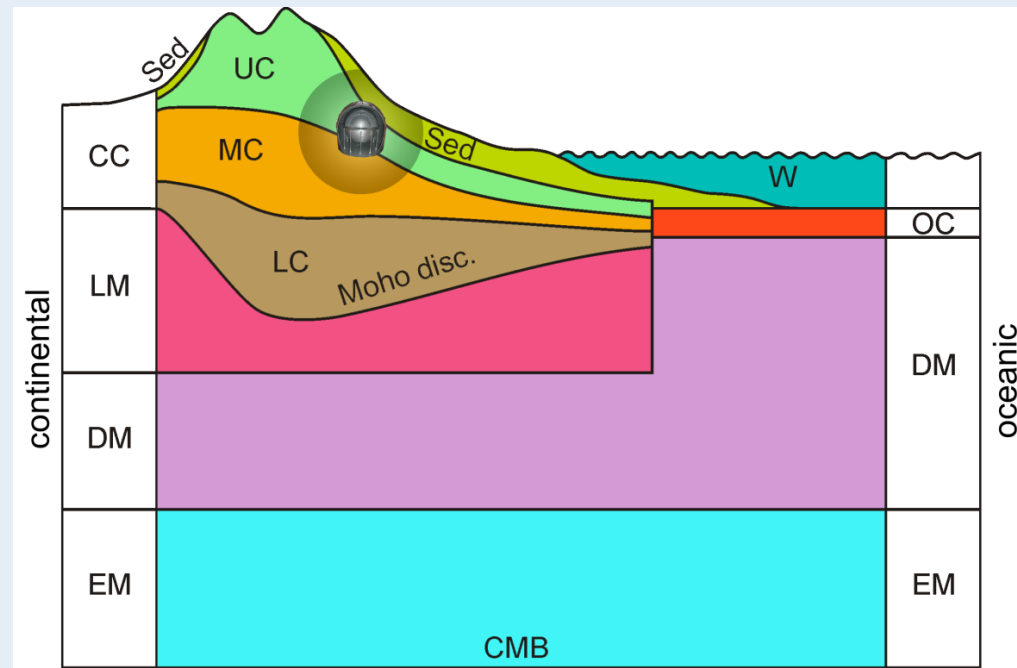
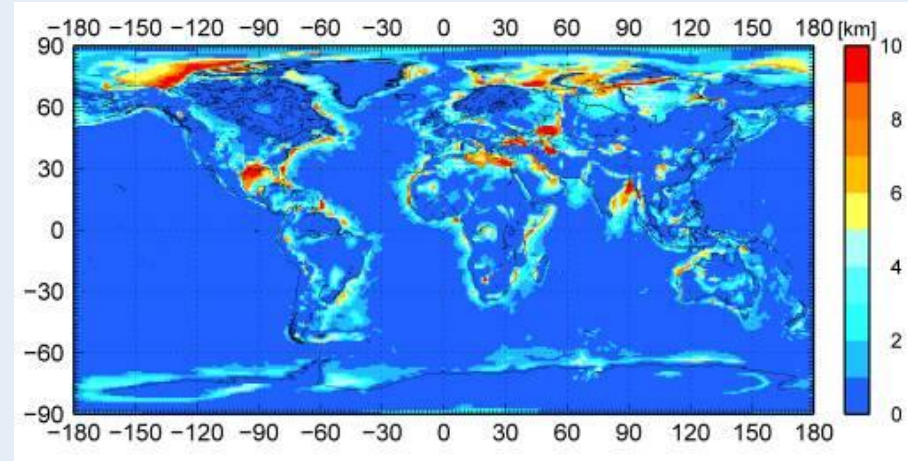




# Structure of RRM and its uncertainties

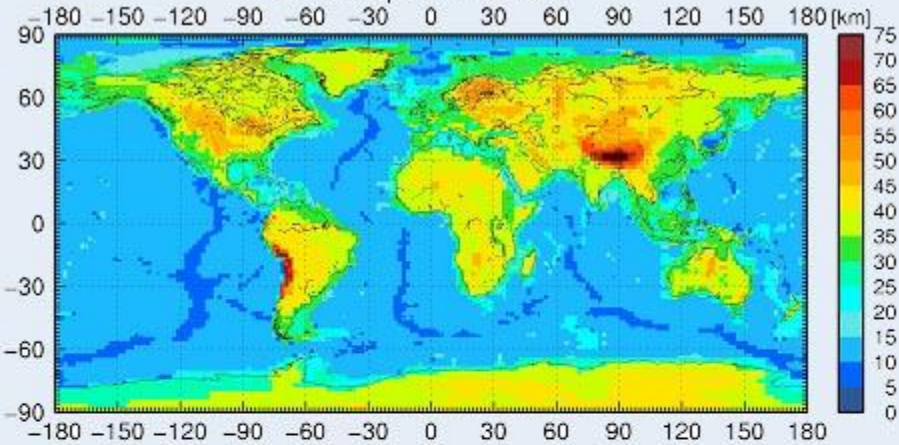
- 1) Topography and bathymetry are adopted from a standard database (ETOPO-5)
- 2) Sedimentary layer - a global  $1^\circ \times 1^\circ$  sediment map [Laske and Masters, 1997]
- 3) Continental crust:
  - a. Upper crust
  - b. Middle crust
  - c. Lower crust
- 4) Oceanic crust
- 5) Lithospheric mantle, depth of CLM base is  $175 \pm 75$  km
- 6) Depleted mantle
- 7) Enriched mantle – 17% of the total mantle mass (Arevalo et al.[2012])

Thickness of sediments ( $1^\circ \times 1^\circ$ )



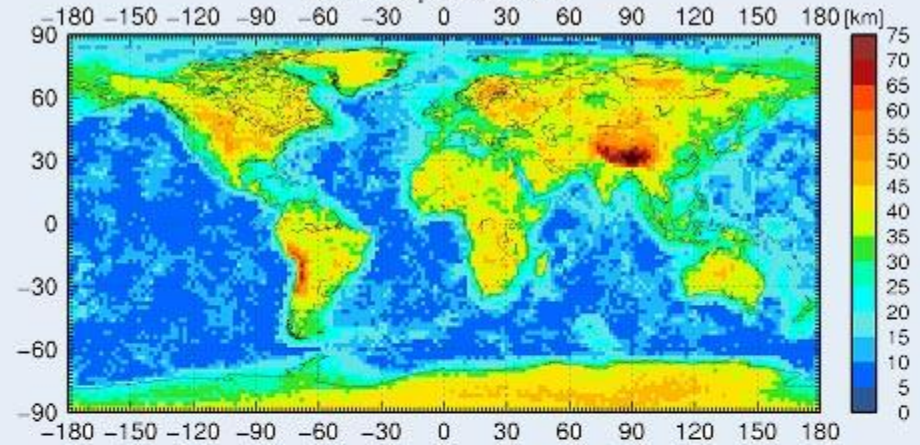
# Crust thickness of RRM

Moho depth CRUST2.0



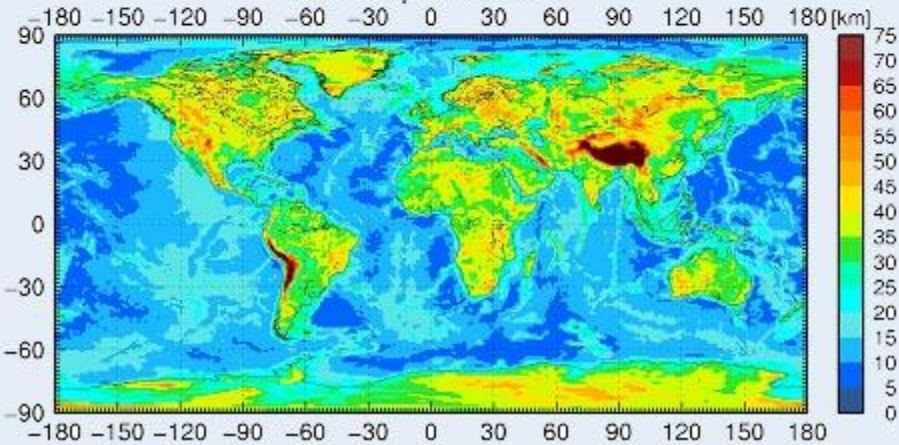
CRUST2.0 (Bassin et al., 2000),  $2^\circ \times 2^\circ$  resolution model based on reflection and refraction seismic body waves

Moho depth CUB2.0



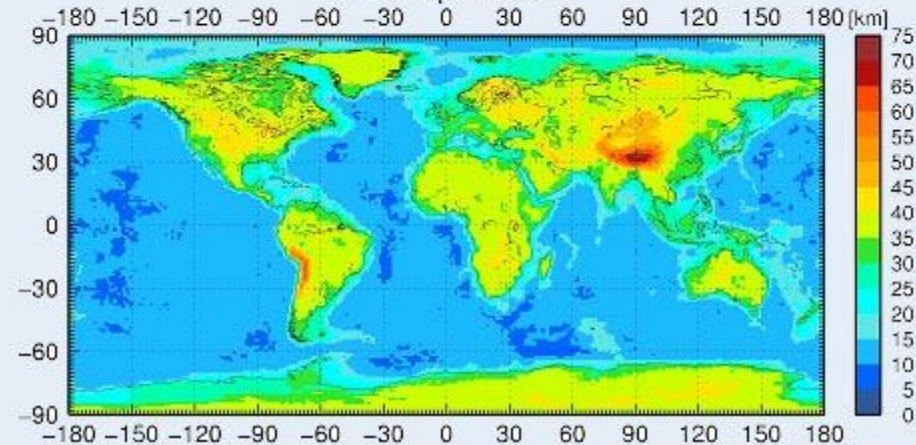
CUB2.0 (Shapiro and Ritzwoller 2002),  $2^\circ \times 2^\circ$  resolution model based on surface wave dispersion

Moho depth GEMMA



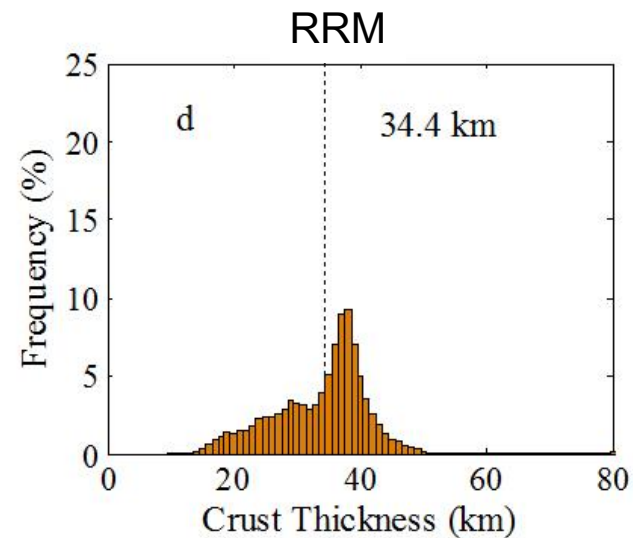
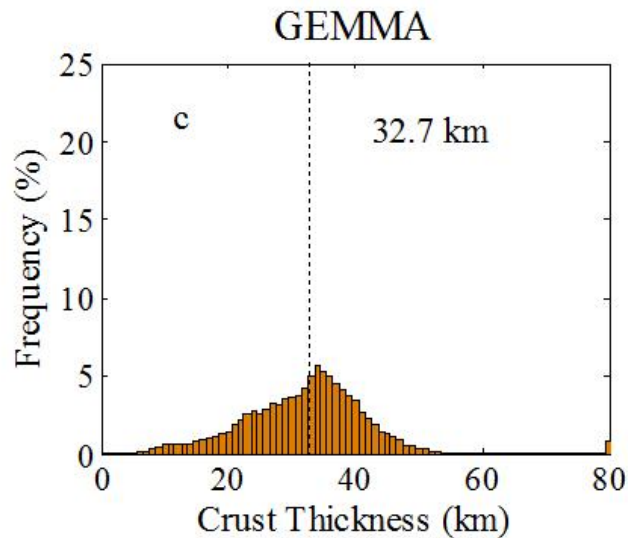
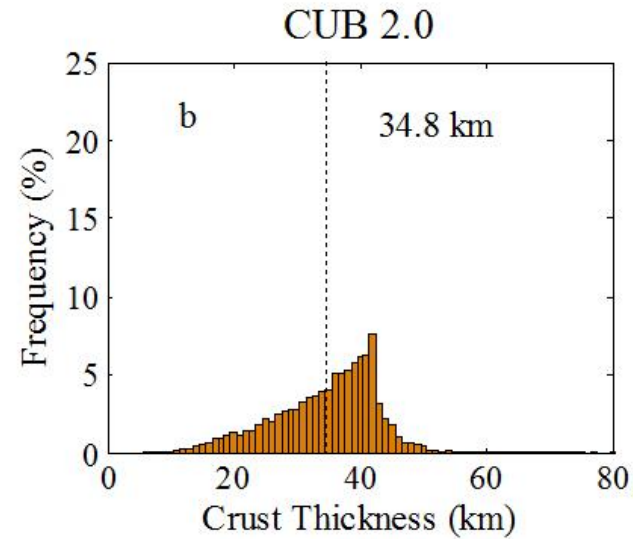
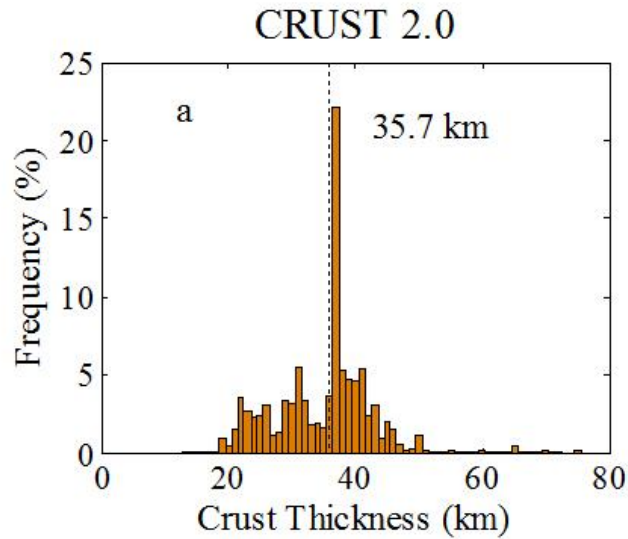
GEMMA (Reguzzoni, Tselfes, 2009; Reguzzoni, Sampietro, 2012) – GOCE Exploitation for Moho Modeling and Applications, the first global high-resolution map ( $0.5^\circ \times 0.5^\circ$ ) of Moho depth Based on gravity field measurements

Moho depth RRM



RRM thickness and its associated uncertainty of each  $1^\circ \times 1^\circ$  crustal voxel is obtained as the mean and the half-range of the three models:

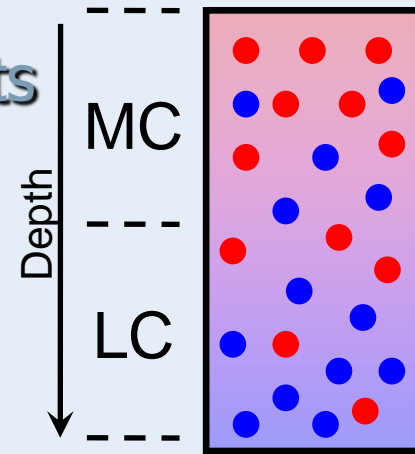
# Distributions of CC thickness in three global crustal models and the RRM



# U and Th in MC and LC : felsic and mafic components

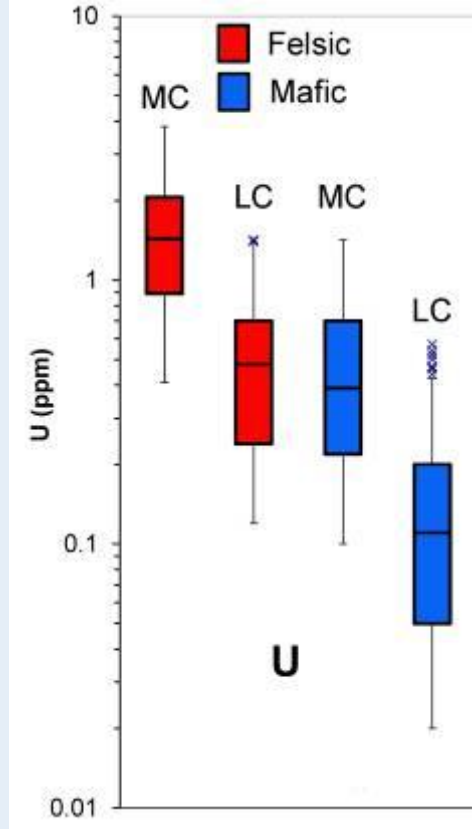
■ In the previous models the abundances of U and Th in MC and LC were affected by large uncertainties, often obtained by average of different compilations.

■ In MC and LC we can recognize two components: **felsic** and **mafic** rocks.



	Fraction	Chemic. Properties	
<b>Felsic</b>	Decrease with depth	Acid	High SiO <sub>2</sub>
<b>Mafic</b>	Increase with depth	Basic	Low SiO <sub>2</sub>

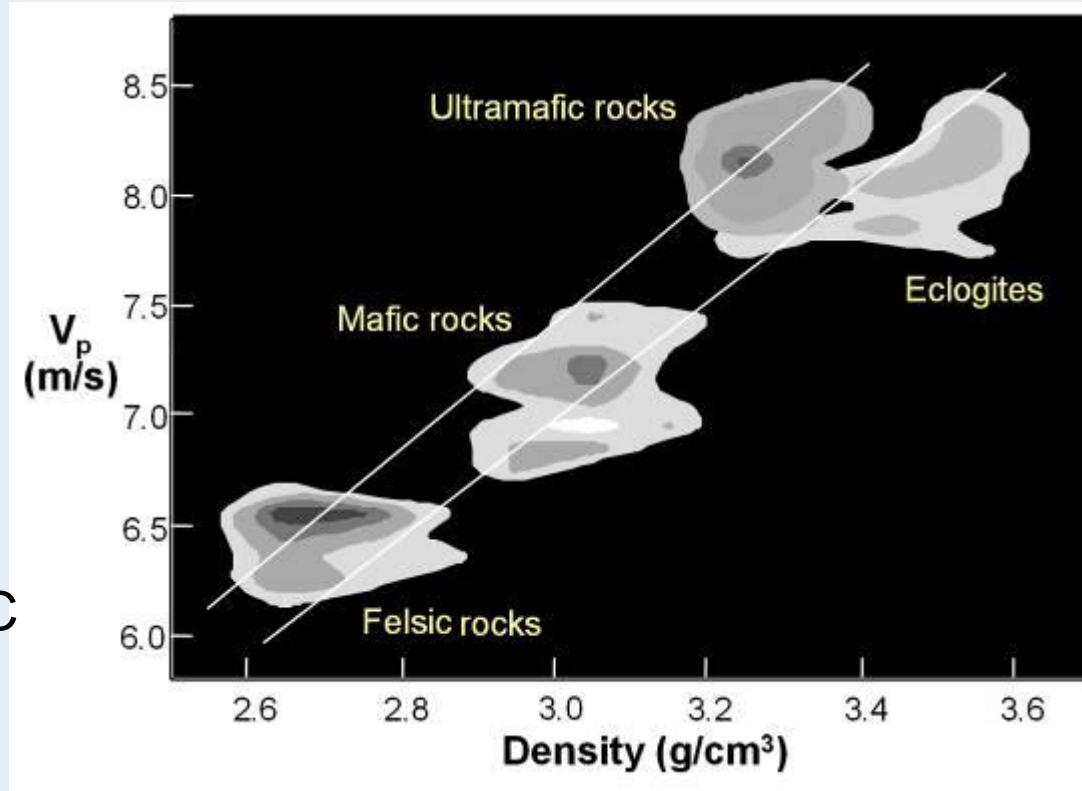
• We compile a database with thousands of published data obtained by ICPMS/ $\gamma$ ray measurements on samples.



		Uranium		Thorium	
		n°	a [ppm]	n°	a [ppm]
MC	<b>Fels.</b>	<b>368</b>	<b>1.4<sub>-0.6</sub><sup>+1.0</sup></b>	<b>428</b>	<b>8.3<sub>-4.1</sub><sup>+8.2</sup></b>
	<b>Mafic</b>	<b>233</b>	<b>0.4<sub>-0.2</sub><sup>+0.4</sup></b>	<b>257</b>	<b>0.6<sub>-0.3</sub><sup>+0.6</sup></b>
LC	<b>Fels.</b>	<b>108</b>	<b>0.4<sub>-0.2</sub><sup>+0.4</sup></b>	<b>133</b>	<b>3.9<sub>-2.5</sub><sup>+7.4</sup></b>
	<b>Mafic</b>	<b>236</b>	<b>0.1<sub>-0.1</sub><sup>+0.1</sup></b>	<b>258</b>	<b>0.3<sub>-0.2</sub><sup>+0.5</sup></b>

# Seismic argument for estimating Felsic/Mafic in MC and LC

- Felsic and mafic rocks can be distinguished on the basis of P and S waves velocities
- Ultrasonic velocity measurements of deep crustal rocks provide a link between seismic velocity and lithology.
- The fractions of **felsic (f)** and **mafic (m)** rocks in the MC and LC of RRM are estimated solving:



$$\begin{cases} f + m = 1 \\ v_C = m v_m + f v_f \\ a = m a_m + f a_f \end{cases}$$

$v_C$  = seismic velocity measured in MC and LC reported in worldwide crustal model CRUST2.0

$v_{f;m}$  = lab. measurements of felsic and mafic rock velocity

$a_{f;m}$  = U (and Th) abundance in felsic and mafic rocks

$a$  = U (and Th) abundance in MC and LC

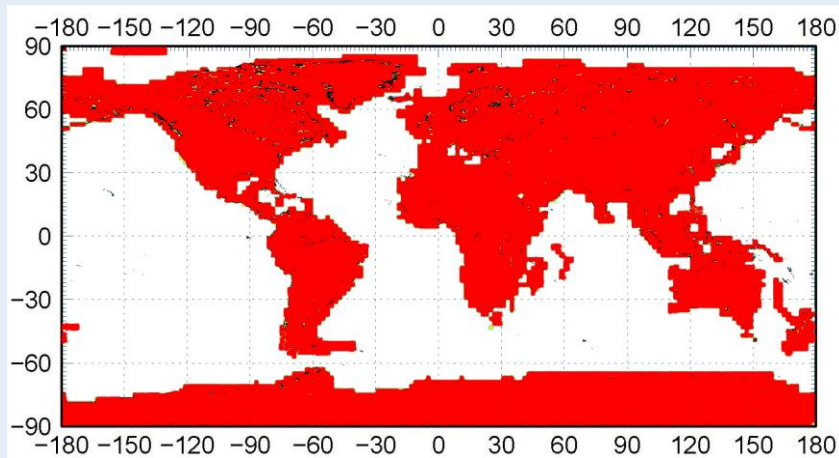
		n° data	$v_p$ (km/s)
MC	Fels.	77	$6.3 \pm 0.2$
	Mafic	57	$7.0 \pm 0.2$
LC	Fels.	29	$6.5 \pm 0.2$
	Mafic	44	$7.2 \pm 0.2$

# U content in the MC and LC: RM vs RRM

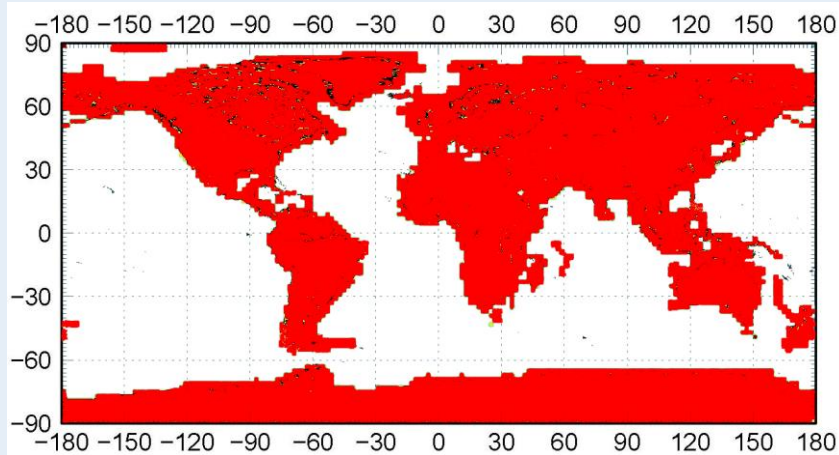
In the RM\* the MC and LC were composed by tiles ( $2^\circ \times 2^\circ$ ) with different thickness but homogeneous content of U: in RRM the U (Th) distribution is heterogeneous

Reference Model\*

U distribution in MC:  $a(U) = 1.6$  ppm

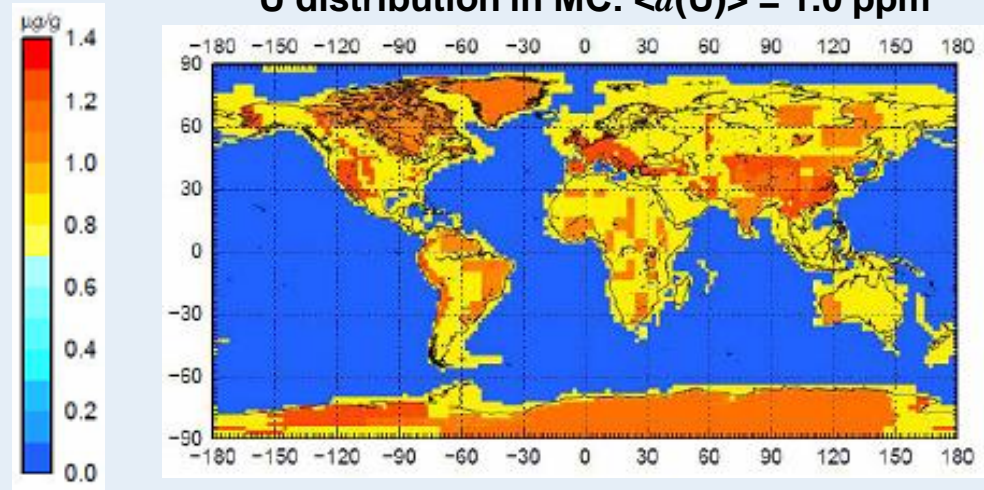


U distribution in LC:  $a(U) = 0.62$  ppm

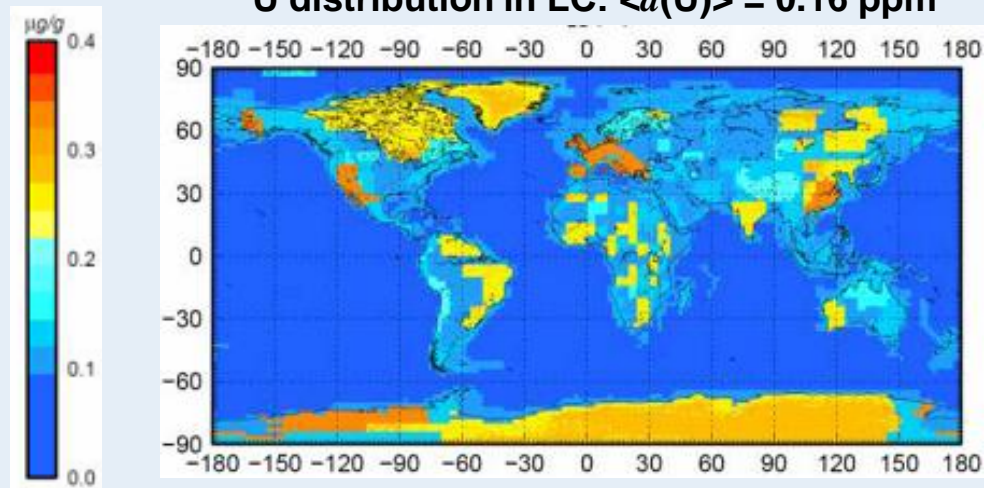


Refined Reference Model\*\*

U distribution in MC:  $\langle a(U) \rangle = 1.0$  ppm



U distribution in LC:  $\langle a(U) \rangle = 0.16$  ppm



\* Mantovani et al. Phys. Rev. D 69, 2004

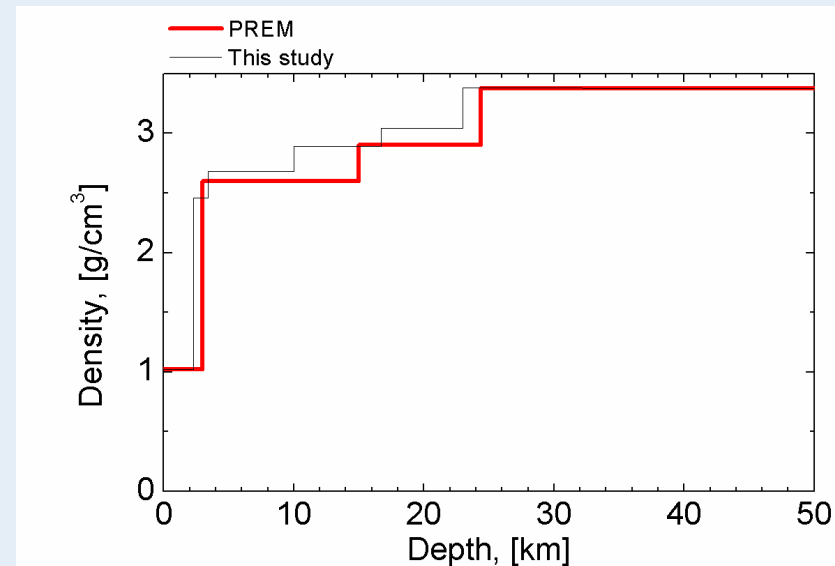
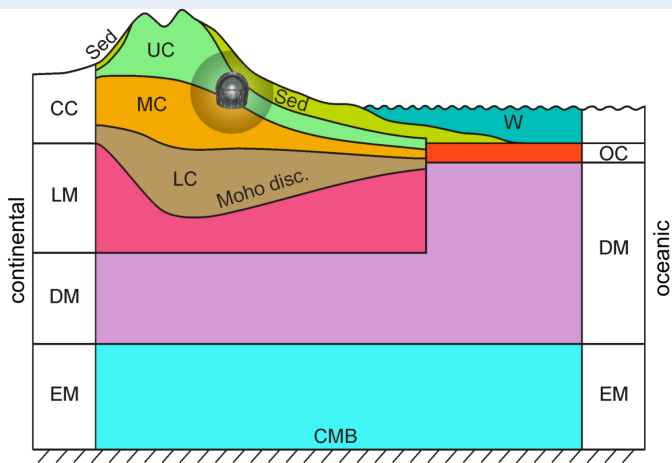
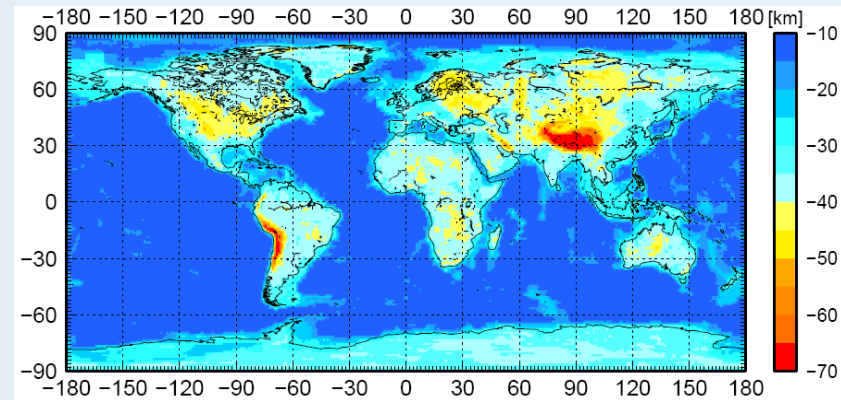
\*\* This results are obtained by using only the  $v_p$  data

# Mantle

1) Heterogeneous topography of the mantle is considered

2) Continental lithospheric mantle is treated as separated reservoir with depth  $175 \pm 75$  km

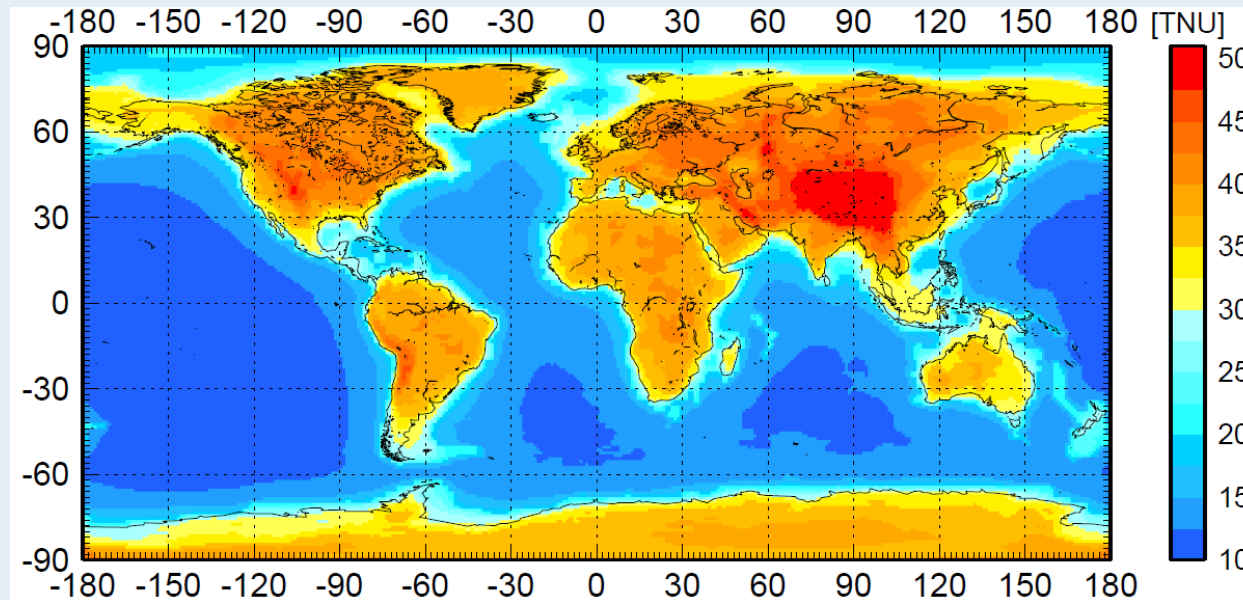
3) Enriched mantle is assumed to contain 17% of the total mantle mass (Arevalo et al.[2012])



# Calculated geoneutrinos signal at LNGS

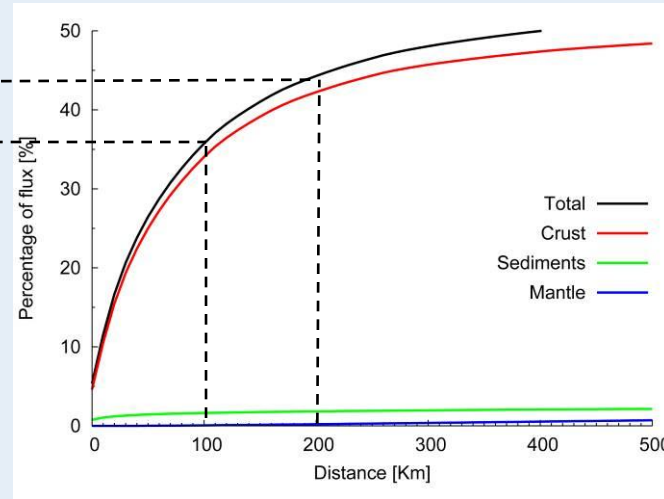
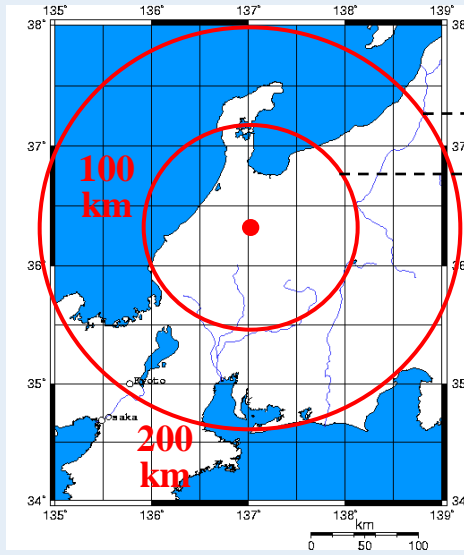
	Borexino (42.45 N, 13.57 E)		
	S(U)	S(Th)	S(U+Th)
Bulk Crust	$21.4_{-4.6}^{+5.2}$	$6.8_{-1.4}^{+2.3}$	$29_{-5.0}^{+6.0}$
CLM	$1.4_{-1.0}^{+2.7}$	$0.4_{-0.3}^{+1.0}$	$2.2_{-1.3}^{+3.1}$
Total LS	$23.6_{-5.2}^{+6.8}$	$7.6_{-1.8}^{+2.9}$	$31.9_{-5.8}^{+7.3}$
DM	4.1	0.8	4.9
EM	2.7	0.8	3.5
Grand Total	$40.2_{-5.8}^{+7.3}$		--

## Geoneutrino signal at Earth's surface



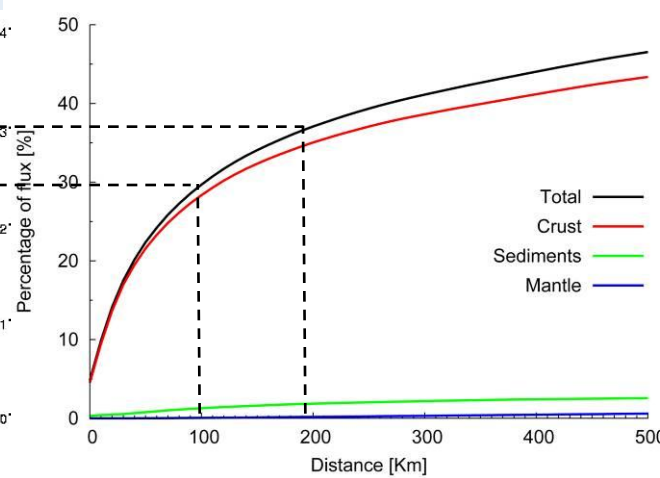
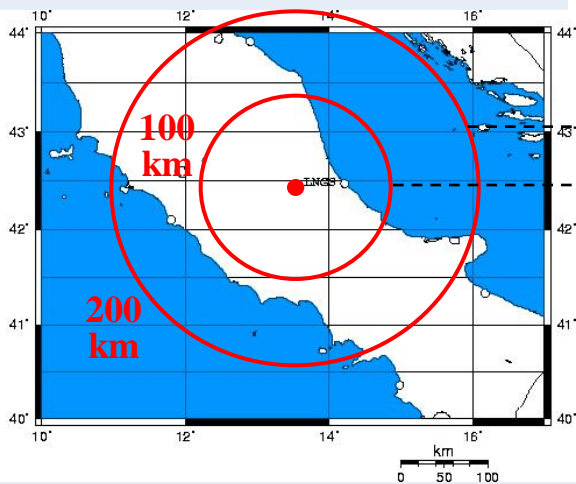


# Why regional geology is relevant?



**KamLAND at Kamioka:**  
50% of the total signal is originated from within 600 km

**Refined Reference Model:**  
*G. Fiorentini et al. – Physical Review D72 – 2005 – arXiv:hep-ph/0501111*



**Borexino at Gran Sasso:**  
• 50% of the total signal is originated from within 900 km

The regional contribution has to be controlled/determined by study of regional geology, if one wants to extract the global information brought in by geo-n's

# Refining of the Reference Model for Gran Sasso

- The reference model [Mantovani et al. 2004] predicts for Borexino:

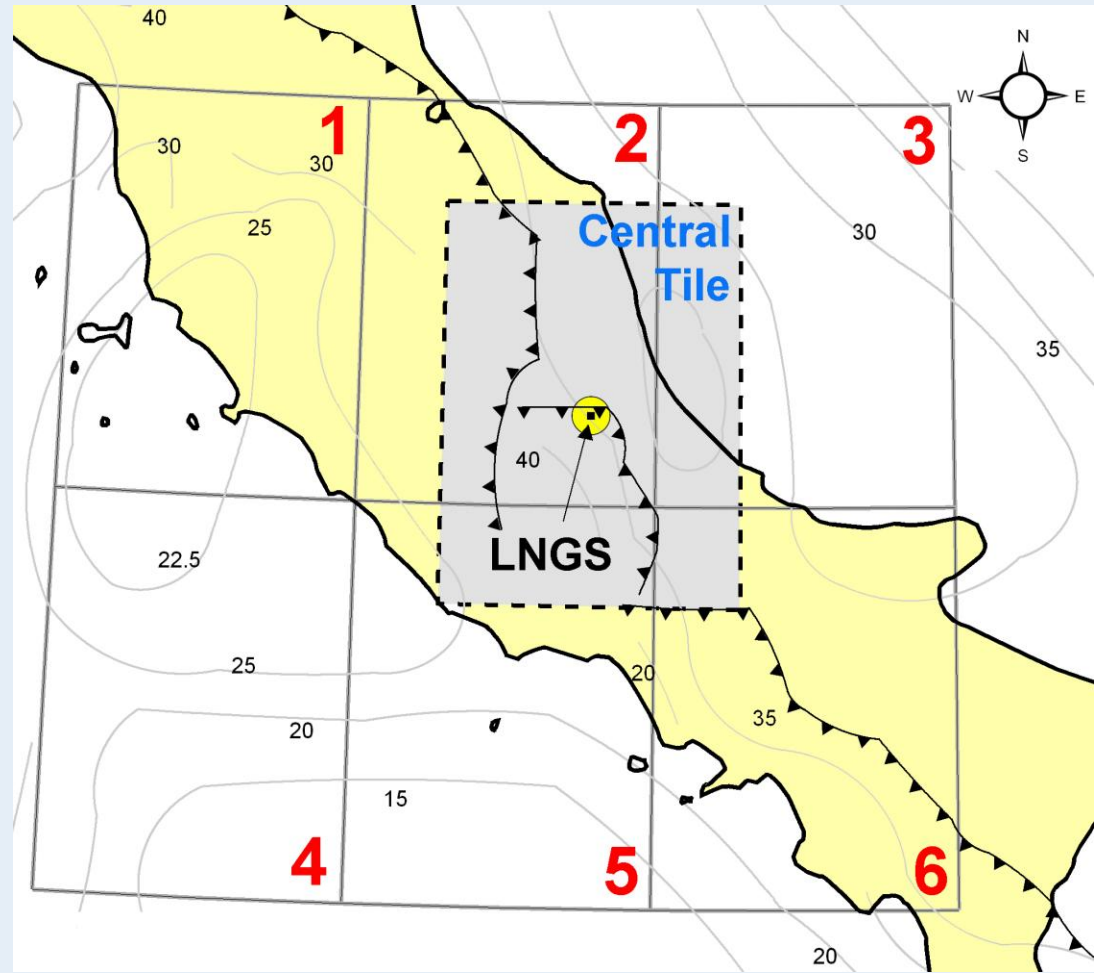
$$S_{\text{Crust}} = 32 \text{ TNU}$$

- The 6  $2^\circ \times 2^\circ$  tiles near Borexino contributes:

$$S_{\text{reg}} = 15.6 \text{ TNU}$$

- A  $2^\circ \times 2^\circ$  tile centered at Gran Sasso gives:

$$S_{\text{CT}} = 12.3 \text{ TNU}$$



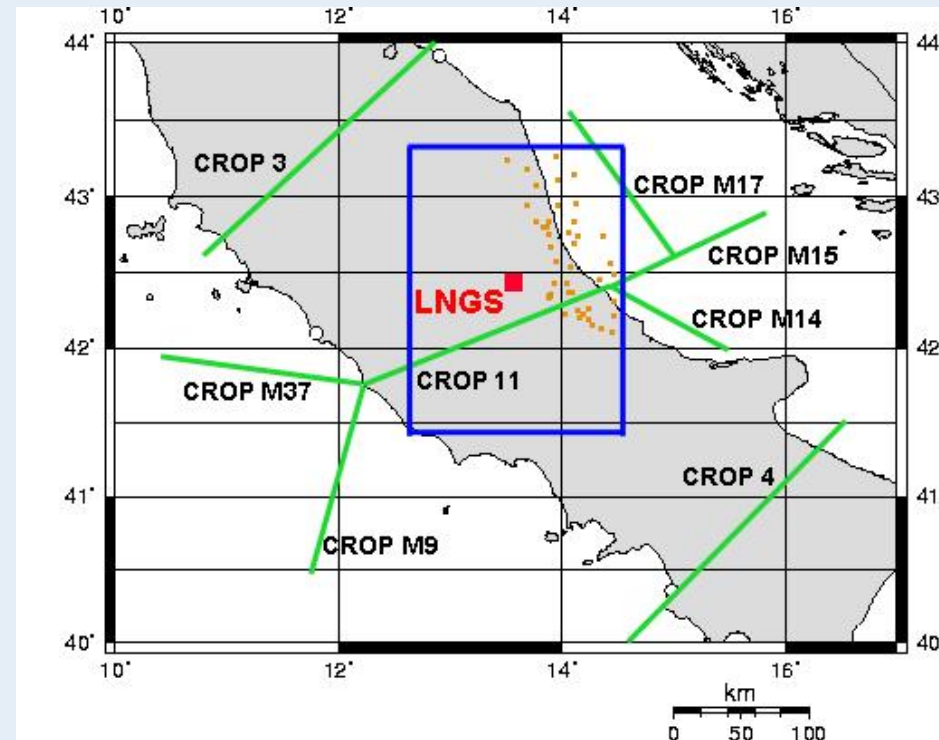
# 3D model of the central tile

## Input:

- Data of **CROP seismic sections**
- Data from 38 **deep oil and gas wells**

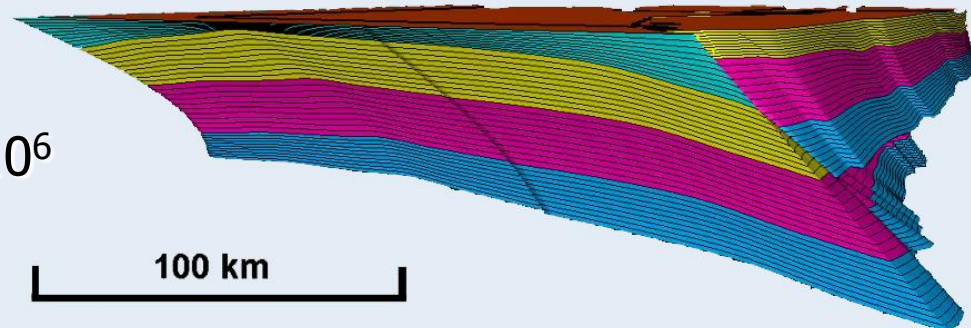
## Identify six reservoirs:

- Sediments:
  - Cenozoic terrigenous units
  - Meso-Cenozoic Basinal Carbonate units
  - Mesozoic Carbonate units
  - Permian and Paleozoic clastic units
- Upper crust
- Lower crust



## Output:

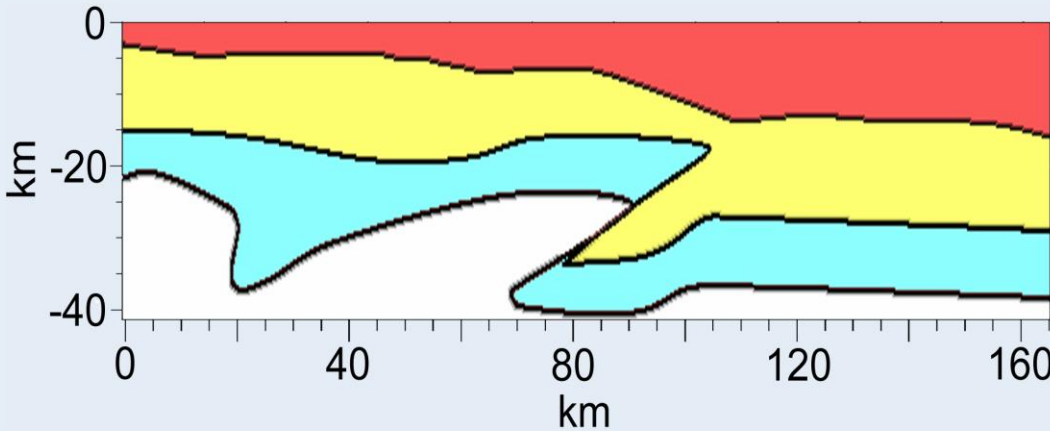
- A 3D dimensional model, built on  $10^6$   $2 \text{ km}^3$  cells



# The sedimentary cover of the central tile

Average thickness of different reservoirs in Reference and 3D model

(the total thickness is almost the same)



Layer	3D model, [km]	Ref. Model, [km]
Sediments	13	0.5
Upper crust	13	10
Middle crust	/	10
Lower Crust	8	10.5
Total	34	31

Measured U and Th content in representative samples of the sedimentary cover:

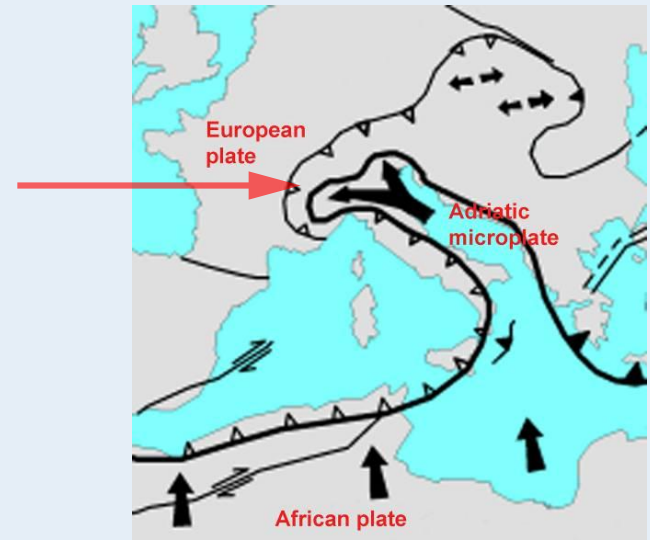
Reservoir	Volume, [%]	$a(U)$ , [ppm]	$a(Th)$ , [ppm]
Mesozoic Carbonate units	74.6	$0.31 \pm 0.19$	$0.21 \pm 0.19$
Cenozoic terrigenous	18.0	$2.28 \pm 0.62$	$8.31 \pm 2.45$
Permian and Paleozoic clastic	5.4	$2.44 \pm 0.70$	$8.76 \pm 2.52$
Meso-Cenozoic Basinal Carbonate	1.9	$2.08 \pm 1.47$	$1.62 \pm 1.76$

# The total signal from the Central Tile

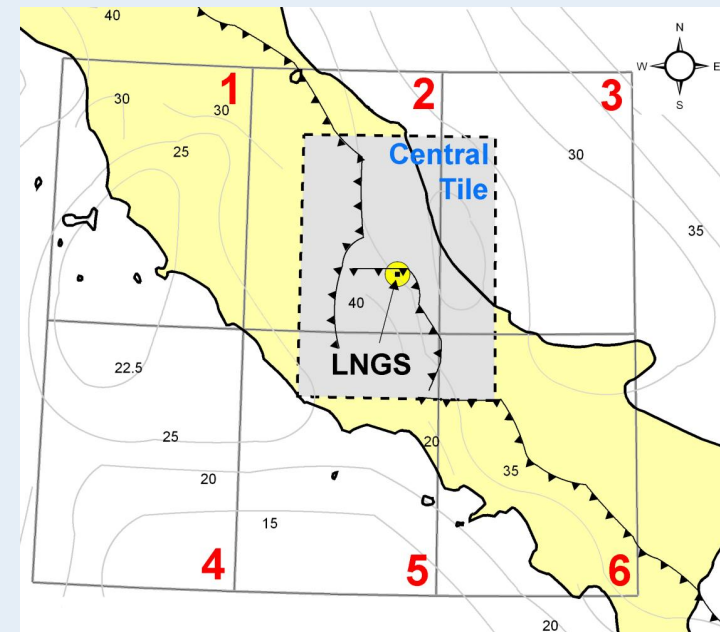
- U and Th content in the Central Tile:

Representative samples from outcropping Adriatic Crust on Western Alps and measured U and Th content

- The total geo-neutrino signal from the Central Tile at Gran Sasso:



Reservoir	Signal from Ref. Model [TNU]	Signal from 3D model [TNU]
Sediments	0.5	2.7
Crust	11.8	5.1
Total	12.3	7.8



# The rest of the regional contribution

We consider the **region out of** the central tile and refine the model by using:

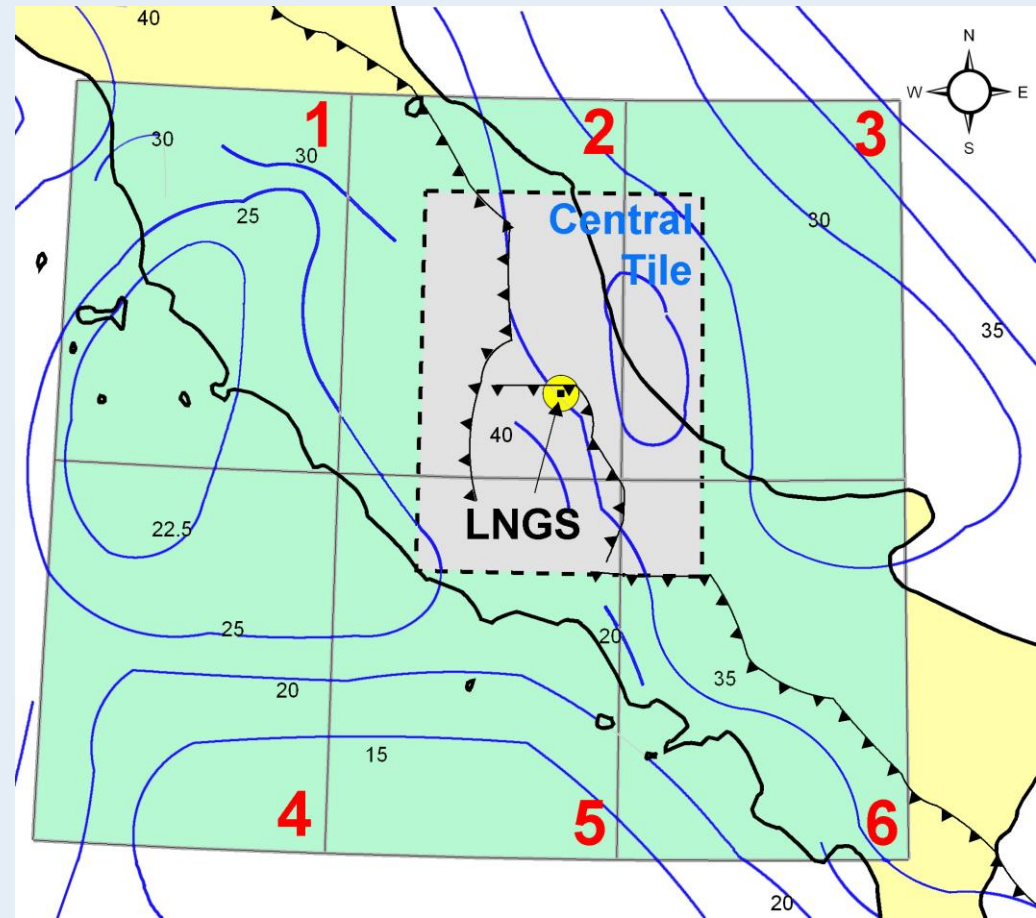
- Data of the main CROP seismic sections
- Depth conversion velocities of the crustal stratigraphic layers
- Detailed measurements of **Moho depth**

By using the deduced U and Th abundance, the signal from the refined **out-region** is:

$$S_{\text{Out}} = 2.3 \text{ TNU}$$

To be compared with that estimated in **Ref.:**

$$S_{\text{Out}} = 3.3 \text{ TNU}$$



# Geoneutrino signal from local geology

Area and reservoir	$S_{RM}(U+Th)$ , [TNU]	$S_{RRM}(U+Th)$ , [TNU]
Central Tile, 3D Model	12.3	7.8 ↓
Rest of the regional area	3.3	2.3 ↓
Regional Contribution, total	15.6	$10.0 \pm 1.3$ ↓

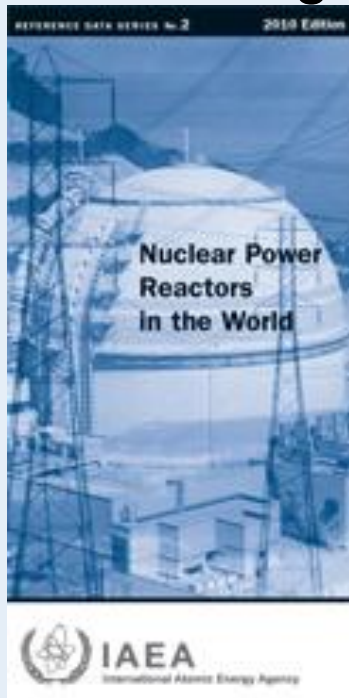
## Total geoneutrino signal at LNGS

$$S_{geo}(\text{Crust}) = S_{geo}(\text{LOC}) + S_{geo}(\text{ROC})$$

	Borexino (42.45 N, 13.57 E)		
	S(U)	S(Th)	S(U+Th)
LOC	$8.1 \pm 1.0$	$1.9 \pm 0.3$	$10.0 \pm 1.3$
ROC	$10.3_{-2.2}^{+2.6}$	$3.2_{-0.7}^{+1.1}$	$13.7_{-2.3}^{+2.8}$
CLM	$1.4_{-1.0}^{+2.7}$	$0.4_{-0.3}^{+1.0}$	$2.2_{-1.3}^{+3.1}$
DM	4.1	0.8	4.9
EM	2.7	0.8	3.5
Grand Total	$34.3_{-2.9}^{+4.4}$		--

# Data Source: IAEA files

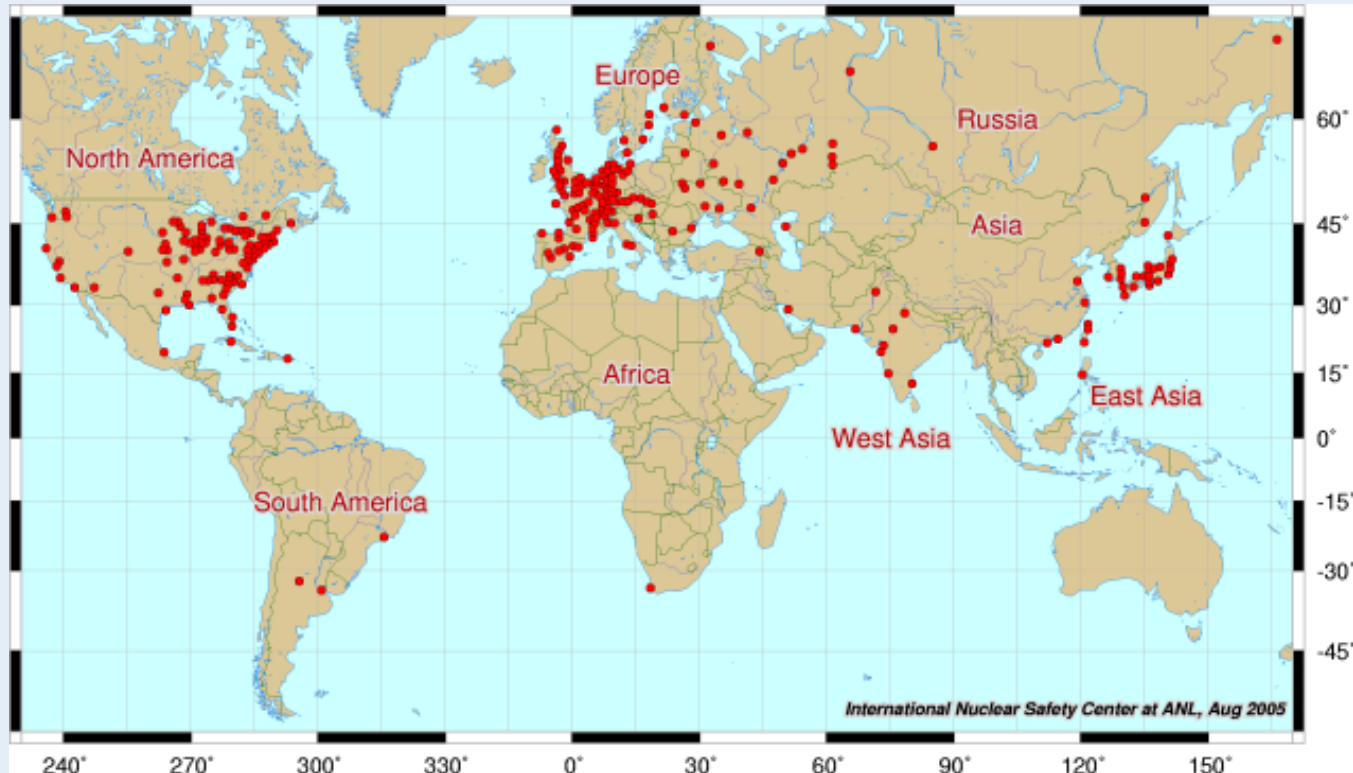
- International Atomic Energy Agency  
<http://www.iaea.org/programmes/a2/>
- On June, description and history of each core are published, referring to previous year.



- Data on: thermal power, electrical capacity, electrical Load Factor, fuel enrichment...



# Nuclear power plants in the world

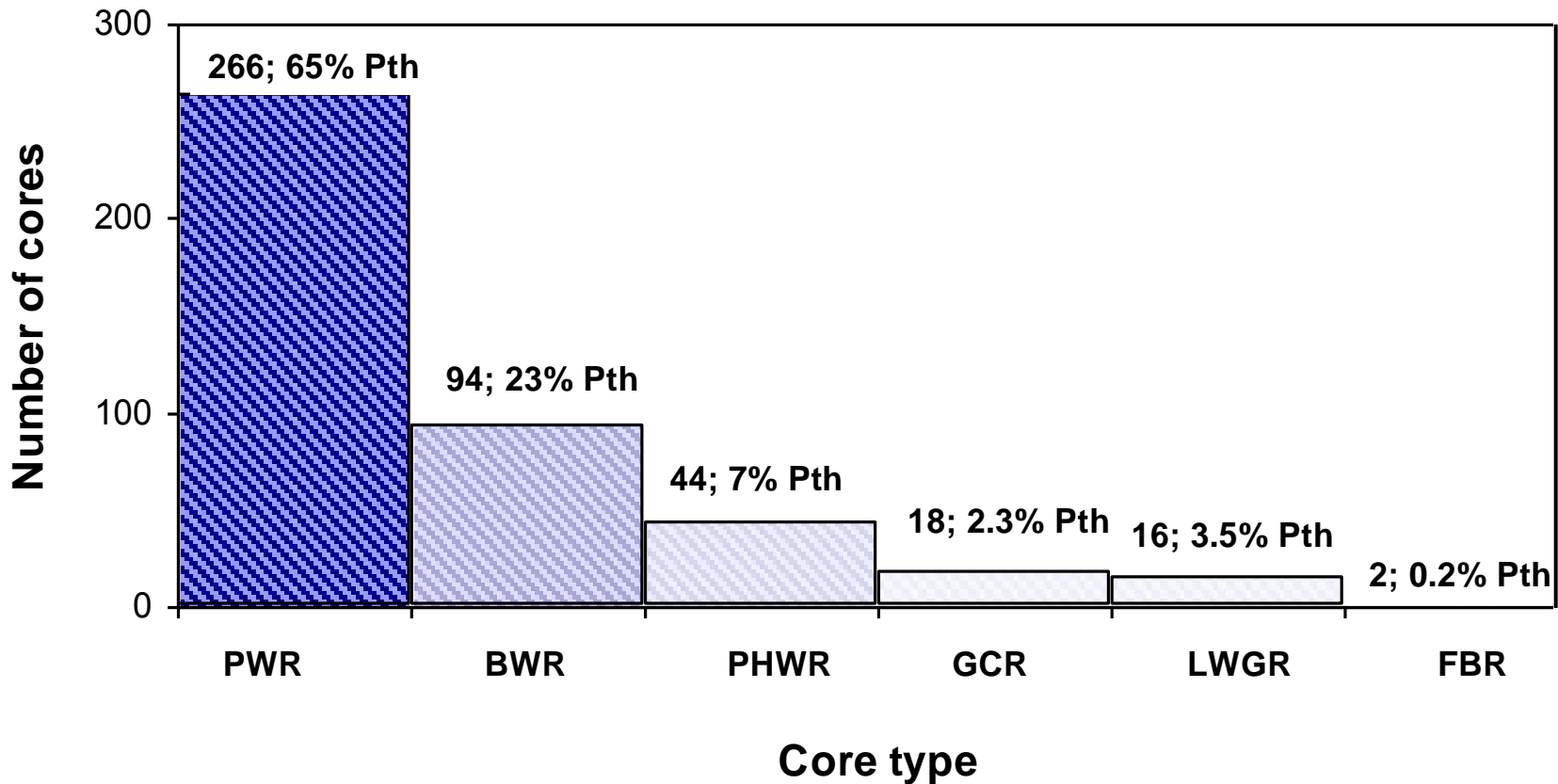


	#cores	$P_{th}$ [GW]
• Europe + Russia	197	519
• North America	122	353
• Japan+ Korea	76	201
• Others	45	75
• Total:	440	1148

- Mean thermal power for core: 2.6 GWth

# Reactors by type



PWR	Pressurized (light) Water Reactor
BWR	Boiling Water Reactor
PHWR	Pressurized Heavy Water Reactor (includes 47 CANDU reactors)

GCR	Gas Cooled Reactor
LWGR	Light Water Graphite mod.
FBR	Fast Breeder Reactor

# Signal calculation

- DETECTOR**
- $\varepsilon=100\%$  detection efficiency
  - $\tau = 1$  year
  - $N_p=10^{32}$

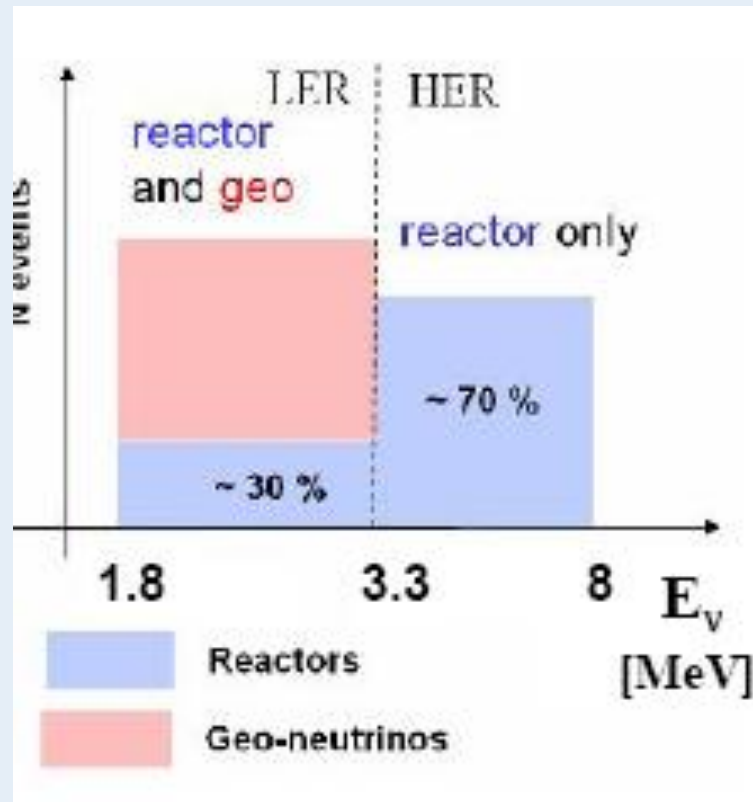
- $\nu$  PHYSICS**
- $P_{ee}$  = survival probability
  - $\sigma(E)$  = cross section  
anti- $\nu_e + p \rightarrow e^+ + n$   
 $E_{th}=1.806$  MeV  
( calculation from Vissani and Strumia 2003)

$$N_{TOT} = \varepsilon N_p \tau \sum_{i=1}^{N_{reactor}} \frac{P_i}{4\pi d_i^2} \langle LF_i \rangle_{2009} \int dE_\nu \sum_{k=1}^{N_{fuel}} \frac{p_k}{Q_k} \lambda_k(E_\nu) P_{ee}(E_\nu, d_i) \sigma(E_\nu)$$

- REACTOR**
- $d_i$  = reactor distance
  - $P_i$  = reference thermal power
  - $LF$  = Load Factor
  - $p_k$  = power fraction

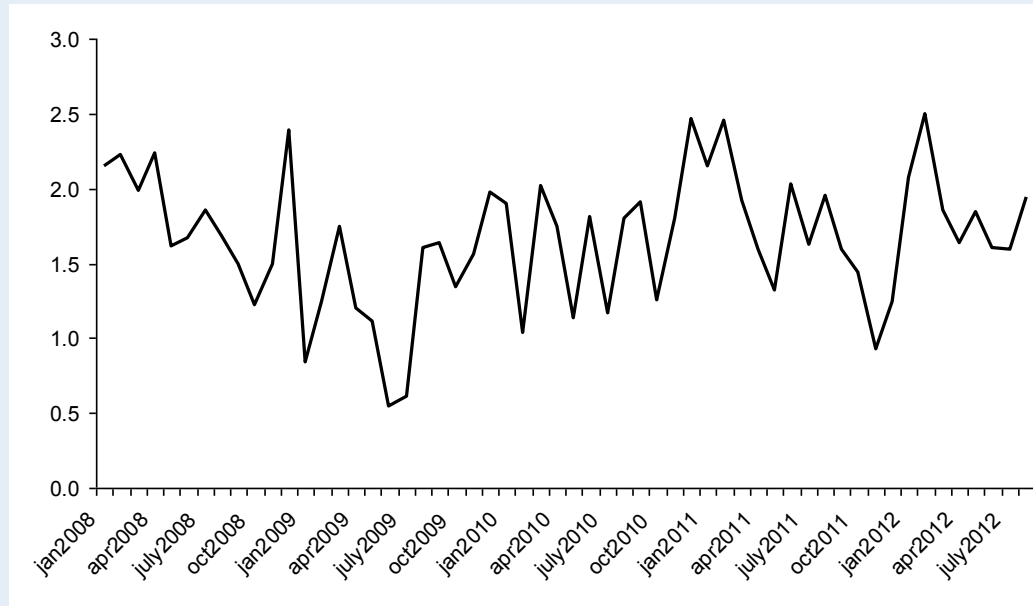
- NUC. PHYS.**
- $Q_k$  = energy released for fission
  - $\lambda_k$  = reactor anti-neutrino spectrum

# Reactor antineutrino signal at LNGS

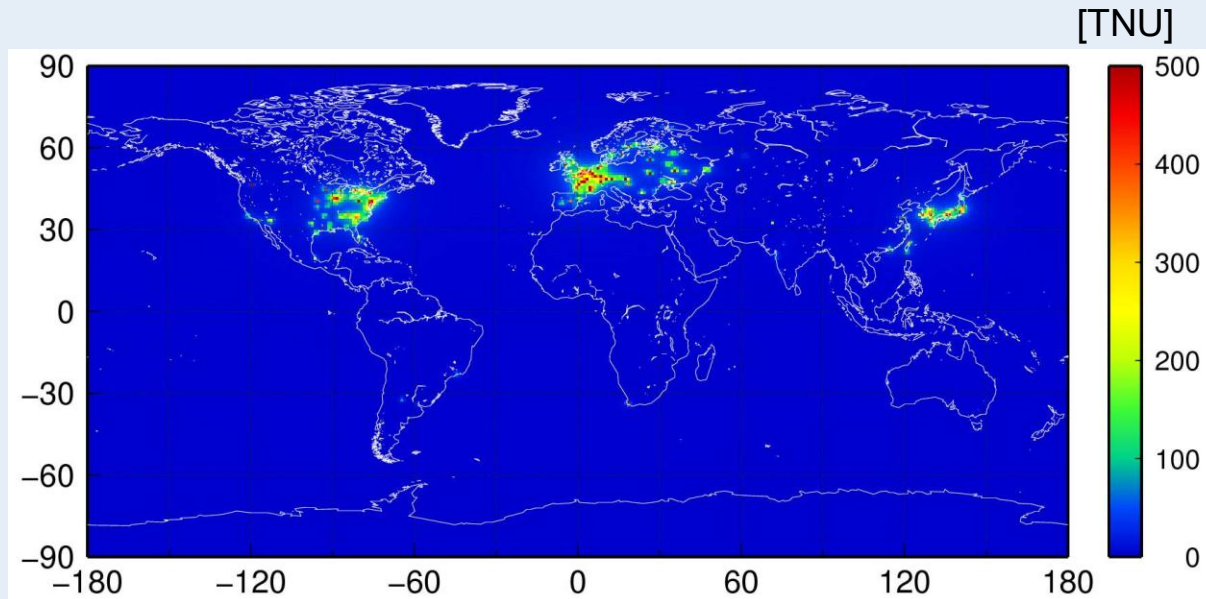


	$S_{\text{react}}, \text{TNU}$
LER	$23.6 \pm 1.2$
HER	$65.2 \pm 3.3$
Total	$88.8 \pm 4.4$

# Monthly evolution of the reactor antineutrino event rate at LNGS



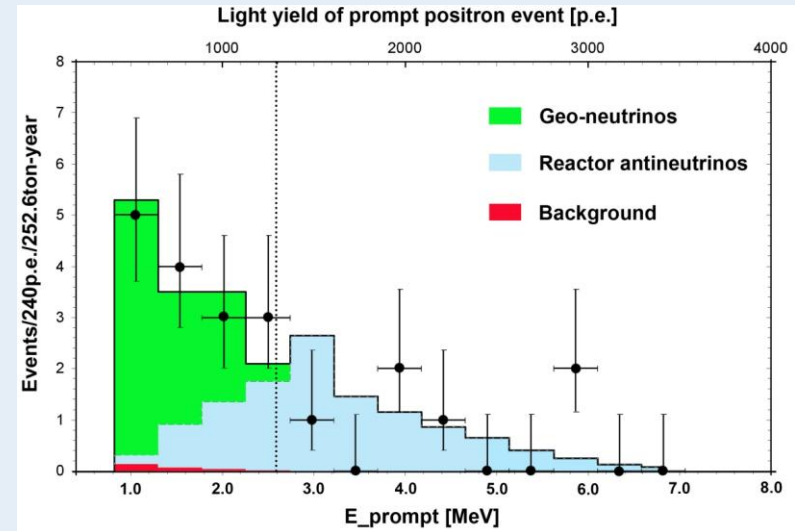
# Reactor antineutrino signal in geoneutrino energy window



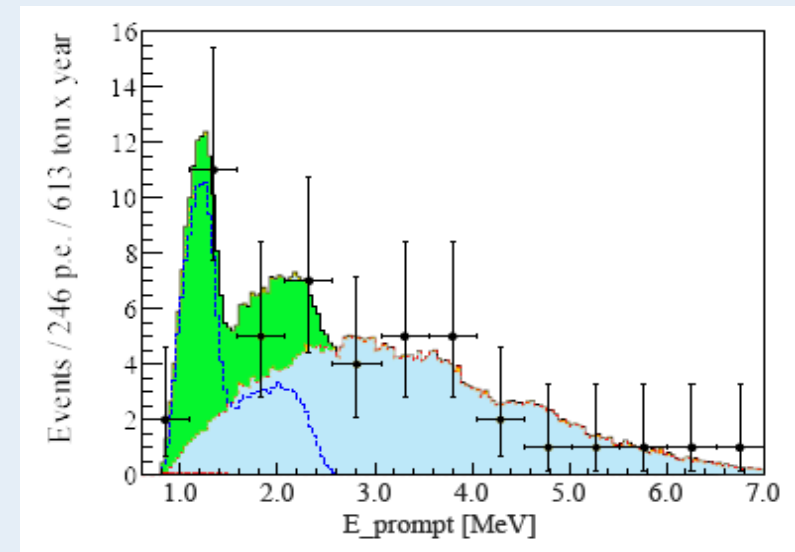
# Geo-neutrino measurements by Borexino



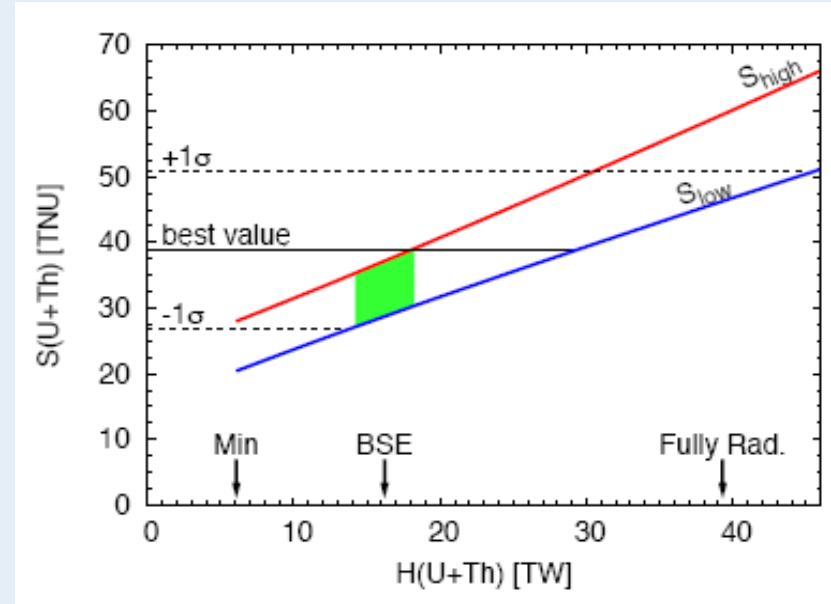
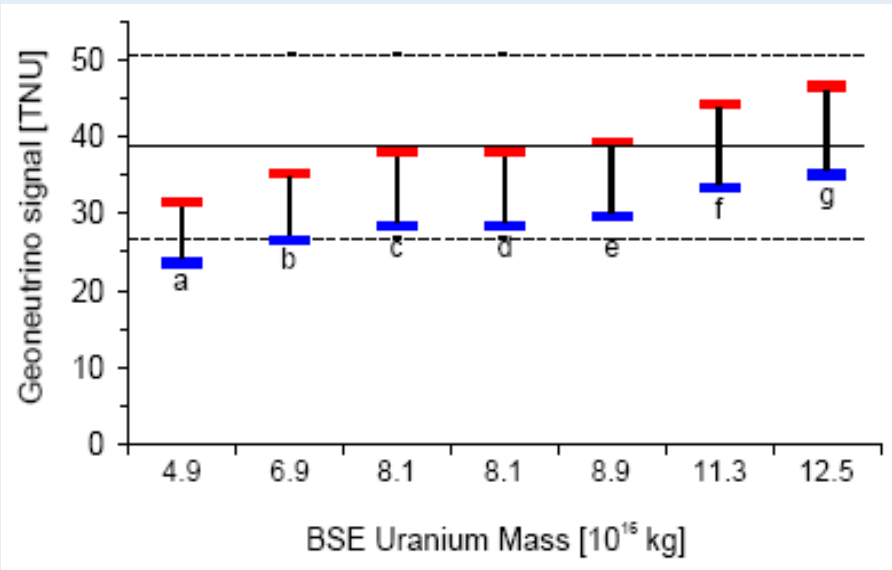
Period	Dec. 07 – Dec. 09	
Exposure p.yr	$0.15 \times 10^{32}$	
	events	TNU
Total (full sp.)	21	
Reactors	$10.7_{-3.4}^{+4.3}$	$64_{-22}^{+26}$
Geo-n	$9.9_{-3.4}^{+4.1}$	$69_{-22}^{+28}$
Background	$0.4_{-0.05}^{+0.05}$	



Period	Dec. 02 – Aug. 12	
Exposure p.yr.	$0.37 \times 10^{32}$	
	events	TNU
Total (full sp.)	46	
Reactors	$31.2_{-6.1}^{+7.0}$	$84.5_{-16.9}^{+19.3}$
Geo-nu	$14.3 \pm 4.4$	$38.8 \pm 12$
Background	$0.70 \pm 0.18$	




# Borexino (2013): geological implications



Geoneutrino signal  $S_{\text{geo}}$  in Borexino (solid line)  $1\sigma$  uncertainty (dashed lines)  $S_{\text{geo}}$ (LOC + ROC+mantle) according to RRM and seven BSE models:

- a) Javoy et al. 2010
- b) Lyubetskaya and Korenaga 2007
- c) McDonough and Sun 1995
- d) Allegre et al. 1995
- e) Palme and O'Neil 2004
- f) Anderson 2007
- g) Turcotte and Schubert 2002

 region allowed by BSE  
 $m(\text{U}) = 0.8 \cdot 10^{17}$  kg  
 [McDonough and Sun, 1995],  
 region between lines contains all  
 models consistent with geochemical  
 and geophysical data

# Conclusions

- 1) With the aim of estimation of the total geoneutrino signal at Borexino and its associated uncertainties the detailed global and local models of crust were developed.
- 2) The geoneutrino signal at Borexino for the new reference Earth model was calculated:  $S_{\text{geo-}\nu} = 34.3^{+4.4}_{-2.9}$  TNU
- 3) For the first time thickness of crust was evaluated from 3 different geological models, abundances in the middle and lower continental crust were calculated by using seismic arguments, continental lithospheric mantle is treated as a separate geochemical reservoir.
- 4) The antineutrino background from nuclear reactors was examined. Operating experience of all commercial nuclear reactors in the last four years was collected. Different power fractions for 3 main types of reactor's fuel burn cycles were considered.



# Written articles

1. Y. Huang, V. Chubakov, F. Mantovani, R. L. Rudnick, W. F. McDonough . A reference Earth model for the heat producing elements and associated geoneutrino flux. *Geochem. Geophys. Geosyst.* 2013.
2. Borexino Collaboration. Lifetime measurements of  $^{214}\text{Po}$  and  $^{212}\text{Po}$  with the CTF liquid scintillator detector at LNGS. *Physical Review C*, 2012.
3. M. Coltorti, R. Boraso, F. Mantovani, M. Morsilli, G. Fiorentini, A. Riva, G. Rusciadelli, R. Tassinari, C. Tomei, G. Di Carlo, V. Chubakov. U and Th content in the Central Apennines continental crust: a contribution to the determination of the geo-neutrinos flux at LNGS. *Geochimica et Cosmochimica Acta* vol. 75, n. 9, 2271-2294, 2011.
4. G. Fiorentini, V. Chubakov, F. Mantovani, B. Ricci. Radiogenic contribution to Earth's heat flow studied through geo-neutrinos. XIV International Workshop on "Neutrino Telescopes". 15-18 March 2011, Venice.
5. G. Fiorentini, M. Lissia, F. Mantovani, V. Chubakov. Geo-neutrinos and radiogenic contribution to Earth's heat flow, AIP Conference Proceedings, 2010.

THANK YOU FOR YOUR ATTENTION

# Participation in the conferences

1. The XXV International Conference on Neutrino Physics and Astrophysics  
3-9 June 2012 (Neutrino) – Kyoto (Japan).  
Title: Reactor antineutrinos in the world.
2. 12th International Conference on Topics in Astroparticle and Underground Physics  
5-9 September 2011 (TAUP) – Munich (Germany).  
Title: Towards a refined reference Earth model for geo-neutrinos.
3. Center for Theoretical Underground Physics and Related Areas  
21 June 2011 (CETUP) – Dakota State University – South Dakota (US).  
Title: Geoneutrinos: Global Crust Model and LNGS Study.
4. Carpathian Summer School of Physics 2010  
20 June – 3 July, 2010, Sinaia, Romania.
5. International Neutrino Summer School 2010  
23-31 August, 2010, Yokohama, Japan
6. 89<sup>o</sup> Congresso SIMP: L'evoluzione del sistema terra dagli atomi ai vulcani  
13-15 September 2010, Ferrara, Italy
7. IV International Neutrino Geoscience 2010 Conference  
6-8 October, 2010, Laboratori Nazionali del Gran Sasso, Italy

THANK YOU FOR YOUR ATTENTION